TO: BOARD OF DIRECTORS

FROM: BRUCE BUEL **DEP**

DATE: MARCH 21, 2008

SAIC PRESENTATIONS

AGENDA ITEM

C-1

MAR 26,2008

ITEM

Dr. Brad Newton of SAIC presentations re Fall Groundwater Storage Volumes and Inter-Basin Transfers [NO ACTION REQUESTED].

BACKGROUND

Dr. Brad Newton of SAIC is scheduled to present the two attached memorandums and will be available to answer questions at the direction of the Board.

RECOMMENDATION

Staff recommends that your Honorable Board receive the presentations and ask questions as appropriate.

ATTACHMENTS

- Fall Groundwater Storage Memorandum
- Inter-Basin Transfer Memorandum

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SCIENCE APPLICATIONS INTERNATIONAL CORPORATION WATER RESOURCES ENGINEERING - *CARPINTERIA*

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- 1 TO: Bruce Buel, General Manager Nipomo Community Services District
- 2 FROM: Joel Degner, Drew Beckwith, Alex Pappas, Brad Newton, Ph.D., P.G.
- 3 RE: Spring and Fall Groundwater in Storage above Mean Sea Level 1975 2007
- 4 DATE: March 20, 2008

5 INTRODUCTION

6 7 8 9 10 11 12 13 14 Groundwater surface elevations (GSE) underlying the Nipomo Mesa are regularly measured at many places (wells) across the mesa. Hydrographs from individual wells provide a temporal record of the GSE measurements at one location. GSE increase and decrease in response to recharge from rainfall, discharge from production, and the balance of natural subsurface flows at that place. Fluctuations in GSE vary across the mesa and make it difficult to assess the regional trend in GSE. Integrating GSE measurements made at one time across the mesa and computing groundwater in storage above mean sea level (GWS) provides a metric to understand the available water for production under the Nipomo Mesa and accounts for disparate trends between wells.

15 16 17 18 19 20 21 22 23 The potential for sea water intrusion to the principal production aquifer is of paramount concern to water supply managers and to all residents who rely on the local groundwater for their supply. The balance between the volume of fresh water under the mesa and the flow of fresh water to the ocean sufficient to prevent sea water intrusion to the principal production aquifer is not currently well understood. However, historical annual GWS estimates for Spring and Fall along with measurements of water quality and GSE collected at the sea water intrusion sentinel wells provide meaningful insights to manage levels of GWS so to guard against permanent degradation that would occur if sea water was to intrude the principal production aquifer.

24 25 26 27 Presented herein are estimates of historical annual variability in GWS from 1975 to 2007 based on groundwater surface elevation measurements collected during Spring and Fall across the Nipomo Mesa. Limited measurements of GSE were available for the years 1982, 1983, 1984, 1994 and 1997, thus precluding a reliable estimate of GWS for those years. **T**

28 RESULTS

29 Overall, the Spring GWS is slightly lower (6,000 AF) in 2007 as compared to 1975, whereas 30 the Fall GWS is much lower (25,000 AF) in 2007 as compared to 1975 (Table 1). Historic GWS 31 ranges from 120,000 AF to 60,000 AF annually during the Spring, and from 95,000 AF to 35,000 32 AF annually during the Fall (Figure 1). The Spring and Fall hydrographs increase and decrease 33 concurrently with decadal periods of wet and dry climatic conditions. Two periods of increased 34 GWS are marked by maximums occurring in 1982 and 2001. One period of decreased GWS is 35 marked by a minimum occurring in 1992. Substantial reductions in the GWS occurred during

Iv:\ ucsd (9103 9235 5935) \tasks\general consultation - 9103\activities\tha11 gres 1975 - 2007\supporting docs\20080320 aminal spring and fall gres 1975 - 2007,docx

SAIC Engineering, Inc. A Subsidiary of Science Applications International Corporation 5464 Carpinteria Ave., Suite K • *Carpinteria, CA 93013* • *Telephone 805/566-6400* • *Facsimile 805/566-6427* To: Bruce Buel Re: GWS Spring and Fall 1975 - 2007 Date: March 20, 2008 Page: 2 of 4

1 2 3 4 the historical periods from 1976 to 1980 and from 1987 to 1996. The recovery from these historical lows in GWS occurred over time scales on the order of five years. The one year change in GWS has been as large as 10,000 AF, and is typically on the order of 5,000 AF or less (Table 1).

5 6 7 8 9 10 11 12 13 14 15 16 17 18 Seasonal difference in GWS has generally increased from 1975 to 2007 (Figure 1). The data is scattered about the linear trend line; nonetheless, the trend is likely real and continues to increase. This observation is may be related to the increase in consumptive use (Technical Memorandum dated January 7, 2008, and presented at Public Workshop dated January 30, 2008). Consumptive use directIy impacts the difference between Spring and Fall estimates because typically no recharge to the GWS occurs during the period from Spring to Fall and the greatest production of water from the principal aquifer occurs from Spring to Fall. However, the Spring GWS increase from Fall GWS is expected to be on average consistent with the recharge from rainfall and thus Spring GWS are expected to parallel the Fall GWS values. The departure from the parallel tracking over the longer multiple year period is not currently understood, and may be in response to increased subsurface flow from the steeper hydraulic gradient occurring with the low Fall GWS. Variations in pumping patterns across the mesa may also contribute to the departure of Spring to Fall GWS values where the GSE measurement of an individual monitoring location may exaggerate or diminish the GWS estimates over time.

19 **METHODOLOGY**

20 21 22 23 24 25 26 27 The annual estimates of Spring and Fall GWS from 1975 to 2007 are based on GSE measurements regularly made by San Luis Obispo County Department of Public Works (SLO F DPW), NCSD, USGS, and Woodlands. The integration of GSE data is accomplished by using computer software to interpolate between measurements and calculate the GWS within the principal production aquifer, to ensure a consistent methodology is used, and to produces a repeatable and therefore comparable metric of the water supply available to the Nipomo Mesa. computer software to interpolate between measurements and calculate the GWS within the
principal production aquifer, to ensure a consistent methodology is used, and to produces a
repeatable and therefore comparable metric precluding a reliable estimate of GWS for those years.

28 29 30 31 32 33 34 The amount of GWS under the Nipomo Mesa Management Area (NMMA) was computed by multiplying the saturated volume above sea level with the aerially weighted specific yield (DWR, 2002), excluding bedrock (Figure 11: Base of Potential Water-Bearing Sediments, presented in the report, Water Resources of the Arroyo Grande - Nipomo Mesa Area [DWR 2002]). The amount of GWS under the NMMA was constrained to the boundary determined in Phase III of the trial. Limited data exist in the additional area included in the Phase V boundary, west of the Phase III boundary.

35 36 37 Data provided by DWR, consisting of well completion reports, lithographic logs, electronic logs, and pump tests, were used to develop an understanding of the hydrogeologic conditions underlying the NMMA. A systematic review of these data pertaining to wells used To: Bruce Buel Re: GWS Spring and Fall 1975 - 2007 Date: March 20, 2008 Page: 3 of 4

1 2 3 4 for storage calculations was conducted in order to verify that each well's screened interval is within the principal production aquifer (Paso Robles Formation). Groundwater surface elevation measurements that do not represent water in the Paso Robles Formation were not included in the calculations.

5 Groundwater Surface Elevation Measurements _ "

6 7 8 9 10 11 12 Groundwater surface elevation data were obtained from the San Luis Obispo County Department of Public Works (SLO DPW), NCSD, USGS, and Woodlands (Table 2). SLO DPW measures GSE in monitoring wells during the spring and the fall of each year. Woodlands and NCSD measures GSE in their monitoring wells monthly. For the years 1975 to 1999, available Department of Public Works (SLO DPW), NCSD, USGS, and Woodlands (Table 2). SLO DPW
measures GSE in monitoring wells during the spring and the fall of each year. Woodlands and
NCSD measures GSE in their monitoring wells mon from wells in the proposed Hydrologic Monitoring Program (HMP) were used to estimate GWS.

13 14 15 16 17 The GSE data was reviewed in combination with well completion reports and historical hydrographic records in order to exclude measurements that do not accurately represent static water levels within the principal production aquifer. Wells that do not access the principal Ine GSE data was reviewed in combination with well completion reports and historical
hydrographic records in order to exclude measurements that do not accurately represent static
water levels within the principal productio within the aquifer were not included in analysis.

18 Groundwater Surface Interpolation .

19 20 21 22 23 24 25 The individual GSE measurements from each year were considered to produce a GSE field by interpolation using the inverse distance weighting (IDW) method. In the IDW method, the GSE field values are computed by from the value, weighted by the distance, of adjacent GSE measurements. The interpolation is based on GSE data alone, and does not incorporate structural geology that mayor may not influence the groundwater surface elevations. In places where a groundwater well has a large areal influence, a small change in GSE can produce a proportionally large change in the estimate of GWS.

26 Groundwater Volume Estimate

27 28 The amount of groundwater in storage under the mesa was estimated for the boundary Groundwater Volume Estimate

The amount of groundwater in storage under the mesa was estimated for the boundary

determined in Phase III of the trial. The GWS was estimated by subtracting both the mean sea 29 level surface (elevation equals zero) and the volume of bedrock above sea level from the 30 saturated volume (Table 1). The bedrock surface elevation is based on Figure 11: Base of 31 Potential Water-Bearing Sediments, presented in the report, Water Resources of the Arroyo 32 Grande - Nipomo Mesa Area (DWR 2002). The bedrock surface elevation was preliminarily 33 verified by reviewing driller reports obtained from DWR. The saturated volume above sea level 34 was multiplied by the specific yield of 11.7% to estimate the recoverable amount of GWS. The 35 specific yield was based on the average weighted specific yield for the Nipomo Mesa 36 Hydrologic Sub-Area (DWR 2002, pg. 86).

To: Bruce Buel Re: GWS Spring and Fall 1975 - 2007 Date: March 20, 2008 Page: 4 of 4

1 **REFERENCES**

2 Department of Water Resources (DWR). 2002. Water Resources of the Arroyo Grande -

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3 Nipomo Mesa Area, Southern District Report.

Copy of document found at www.NoNewWipTax.com

Spring and Fall Groundwater in Storage Above Mean Sea Level for Phase III Boundary

--- Insufficient data for evaluation

Figure 1

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SCIENCE APPLICATIONS INTERNATIONAL CORPORATION WATER RESOURCES ENGINEERING - *CARPINTERIA*

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TECHNICAL MEMORANDUM

- 2 TO: Mr. Bruce Buel, General Manager NCSD
- 3 FROM: Drew Beckwith, Brad Newton Ph.D. P.G., Robert G. Beeby, P.E.
- 4 RE: Estimate of 2007 Subsurface Flow Across the Phase III Boundary FROM: Drew Beckwith, Brad Newton Ph.D. P.G., Robert G. Beeby, P.E.
RE: Estimate of 2007 Subsurface Flow Across the Phase III Boundary
DATE: March 11, 2008
INTRODUCTION
- 5

6 INTRODUCTION

7 8 9 10 11 12 13 The historic low rainfall of 2007 (less than 7 inches or less than 50% of normal) resulted in a substantial decrease in groundwater surface elevations (GSE) and a decrease in the groundwater in storage under the Nipomo Mesa (SAlC technical memorandum dated August 28, 2007). This decrease in GSEs underlying the Nipomo Mesa potentially altered hydraulic gradients along the Phase III boundary (henceforth Boundary) and may have affected subsurface flow across specific flow zones as well as net subsurface flow, as compared to previous years. R

14 15 16 17 18 19 20 21 SAlC estimated subsurface flow across the Boundary in year 2007. Two estimates of subsurface flow are presented herein that differentiate between subsurface flow related to fresh water, and a portion of fresh water available for water supply. The difference depends on the definition of the base of the flow zone: 1) Fresh Water was defined following work conducted by Slade and Associates titled, "Map of 'Base of Fresh Water' " dated March 2000 (Slade 2000), and 2) Water Supply was defined as the higher of the data collected by the California Department of Water Resources (DWR) presented in "Base of the Potentially Water-Bearing Sediments" (Plate 11, DWR 2002), or sea level.

22 23 24 25 26 Geologic characteristics (stratigraphy and structure) play an important role in the flow of groundwater. The estimates presented herein are based on the assumption that the aquifer is unconfined and no barriers to flow exist internal to the Boundary. Refinements to the understanding of the local geology and its effect on groundwater flow will occur as data become available and evaluations are conducted. F T

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28 RESULTS

29 Net subsurface flow in year 2007 across the Boundary was 5,050 acre-feet (AF) for Fresh 30 Water, and 1,370 AF for Water Supply. The Boundary was divided into five separate flow 31 zones adjacent to the following areas: Nipomo Valley, Los Berros Creek, Arroyo Grande Plain, 32 Pacific Ocean, and Santa Maria River Valley to estimate the net subsurface flow to Nipomo 33 Mesa (Figure 1). The calculation method is based on Darcy's Law and GSE measurements taken 34 during Spring 2007 (Table 1). The daily estimate of flow from Spring 2007 GSE measurements 35 was converted to its annual equivalent for the purposes of comparison to tangible and

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SAIC Engineering, Inc. A *Subsidiary of Science Applications International Corporation 5464 Carpinteria Ave., Suite K* • *Carpinteria,* CA *93013* • *Telephone 805/566-6400* • *Facsimile 805/566-6427* To: files Re: Subsurface Flow Date: March 10,2008 Page: 2 of 6

1 2 3 4 meaningful values of groundwater in storage above sea level (GWS). The sign convention established herein presents inflow to the Nipomo Mesa as positive values and outflow from the Nipomo Mesa as negative values. Net subsurface flow compares inflows to outflows estimated across the five flow zones.

5 6 7 8 $\dot{9}$ Fresh Water flow is more than three times Water Supply flow because a significant amount of water flows across the Boundary below sea level and therefore is not currently considered as water supply for production. For example, the Fresh Water inflow across the Santa Maria Valley flow zone considered flow through a saturated height on the order of 1,500 feet (ft) (Col. 4, Table 3), as compared to the Water Supply flow which considered flow through

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10 a saturated height on the order of 100 ft (Col. 5, Table 4).

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12 Table 1. Subsurface flow across the Boundary (presented as an annual total) R

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15 16 17 18 Total 1,400 30 1,370
METHODOLOGY
Darcy's Law was used to estimate year 2007 subsurface flow across the Boundary. Darcy's Law describes the flow of a fluid through a porous medium, and states that the flow of water through a cross-section is equal to the product of the hydraulic conductivity, hydraulic gradient, and cross-sectional area, as follows:

$$
Q=K\ast I\ast A
$$

19 where,

- 20 $Q =$ subsurface groundwater flow [cubic feet per day (ft³pd)]; 21 K = hydraulic conductivity [cubic feet per day per square foot (ft³pd/ft²)]; 22 $I = \frac{1}{2}$ = hydraulic gradient [feet per feet (ft/ft)];
- 23 $A = \text{cross-sectional area}$ [square feet (ft²)].

To: files Re: Subsurface Flow Date: March 10, 2008 Page: 3 of 6

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2 3 4 5 6 7 8 9 10 11 12 13 14 15 16 Darcy's Law is straightforward and therefore frequently used to estimate groundwater flow. However, because of data limitations, determining the value of each variable requires some interpretation and can lead to various approaches used to determine the quantity of each barcy s Law is straightforward and therefore riequently used to estimate groundwater
flow. However, because of data limitations, determining the value of each variable requires
some interpretation and can lead to various a estimate of flow. Here, GSE measurements were collected and compiled. Contours of GSE were created on a map by linear interpolation between measurements (Figure 2). Each flow zone was divided into segments (Col. 1 Tables 3 and 4). The flow zones were segmented along the Boundary to account for spatial variations in the elevation of the bottom of the flow zone according to previous work performed by SAIC (SAIC 2002). Cross-sectional area (Col. 5 Table 3, and Col. 6 Table 4) was computed for each segment along each flow zone by multiplying the segment length (Col. 1 Tables 3 and 4) with the saturated height (Col. 4 Table 3, and Col. 5 Table 4) determined from the contour map and base of flow. Hydraulic gradients were measured on the map normal to GSE contours at each segment along each flow zone (Col. 6 Table 3, and Col. 7 Table 4). Hydraulic conductivity was measured and compiled by DWR and attributed to the 4) determined from the contour map and base of flow. Trydiating gradients were measured on
the map normal to GSE contours at each segment along each flow zone (Col. 6 Table 3, and Col.
7 Table 4). Hydraulic conductivity wa

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18 *Groundwater Contour Map*

19 20 21 22 23 Spring 2007 GSE measurements provided by the San Luis Obispo County Department of Public Works were used to create a contour map of the GSE underlying the Nipomo Mesa Spring 2007 GSE measurements provided by the San Luis Obispo County Department of
Public Works were used to create a contour map of the GSE underlying the Nipomo Mesa
(Figure 2). Groundwater surface elevations were assumed measurement locations. Geologic structures were assumed to not influence the groundwater contour shape.

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25 *Cross-Sectiona 1 Area*

26 27 28 29 30 31 Cross-sectional area was defined as the length of the flow zone times the saturated height of the flow. In the Fresh Water estimate, the saturated height was determined as the difference between the elevations from a GSE contour map (Col. 3 Table 3), and elevations from a contour map of the Base of Fresh Water (Col. 2 Table 3). The depth of flow used in this approach considered as much as 1,500 ft of fresh water aquifer in a few segments, consistent with the work previously presented during the Santa Maria Groundwater Litigation.

32 33 34 35 In the Water Supply estimate, the saturated height was determined as the difference between the elevations from a GSE contour map (Col. 4 Table 4), and the higher of 1) the base of the potentially water-bearing sediments (Col. 2 Table 4), or 2) sea level (Col. 3 Table 4). The depth of flow used in this approach considered only the upper portion of the fresh water To: files Re: Subsurface Flow Date: March 10, 2008 Page: 4 of 6

1 aquifer and generally represents the flow that contributes to GWS. In this case, the saturated 2 height is at most 120 ft of the upper portion of the fresh water aquifer in all segments.

3 4 5 The length of each segment, multiplied by the average of the saturated height at each endpoint of the segment defines the cross-sectional area of the segment used in subsurface flow calculations (Col. 5 Table 3, and Col. 6 Table 4).

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7 *Hydraulic Gradient*

8 9 10 11 12 The gradient within each segment is the average of the gradients measured at each segment endpoint from GSE contours at the Boundary (Col. 6 Table 3, and Col. 7 Table 4). D Gradient measurements were made over a length scale of approximately one and a half miles and represent the average of locally steep and shallow variations in GSE present along the Boundary.

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14 *Hydraulic Conductivity*

15 16 17 18 19 20 21 22 Hydraulic conductivity values for stratigraphy underlying the Nipomo Mesa are presented in "Water Resources of the Arroyo Grande-Nipomo Mesa Area" (DWR 2002). SAIC obtained the geometric mean of hydraulic conductivity along segments (flow zones) of the DWR study boundary (personal communication, Evelyn Tompkins, DWR). For each flow zone, the geometric mean of hydraulic conductivity was used in the subsurface flow calculation for all botained the geometric mean of hydraulic conductivity along segments (flow zones) or the LAD
DWR study boundary (personal communication, Evelyn Tompkins, DWR). For each flow zone,
the geometric mean of hydraulic conductivi mean of hydraulic conductivity for the Nipomo Valley flow zone was used (Table 2, Col. 7 Table 3, Col. 8 Table 4).

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24 Table 2. Hydraulic Conductivity Values

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26 DISCUSSION

27 Subsurface flow above sea level relates to GWS available as water supply to the Nipomo 28 Mesa. Long-term production of groundwater below sea level in excess of the recharge will 29 cause groundwater surface elevations underlying the Nipomo Mesa to drop below sea level and To: files Re: Subsurface Flow Date: March 10,2008 Page: 5 of 6

1 2 3 4 induce sea water intrusion. Consequently, the estimate of net subsurface flow presented herein related to the water supply to the Nipomo Mesa is 1,370 AF. It is based on GSE measurements made during Spring 2007, presented as an annual value, and is predicated on the assumptions of an unconfined aquifer with no geologic structural controls.

5 6 7 8 9 10 Deep groundwater flow may be an important component of the hydrologic balance under the Nipomo Mesa, particularly where fresh water flow crossing the Boundary at the Santa Maria Valley and Ocean flow zones is influenced by the hydrostatic pressures of water above sea level and perhaps from water at significant distances from the Boundary. Given the same assumptions, the estimate of net subsurface flow presented herein representing deep groundwater is 5,050 AF.

11 12 13 14 15 16 17 While GSE measurements are compiled on a frequent basis within the Boundary, much fewer measurements have been compiled for the area outside of the Boundary. This data limitation makes GSE contouring difficult, especially near the edges of the Boundary where resultant hydraulic gradient calculations are vital to determining subsurface flow. The Santa Maria Valley flow zone is paramount to the water supply of the Nipomo Mesa. The data miniation makes GSE contouring difficult, especially hear the edges of the boundary where
resultant hydraulic gradient calculations are vital to determining subsurface flow. The Santa
Maria Valley flow zone is paramount to greatest uncertainty.

18 19 20 21 22 Geologic characteristics (stratigraphy and structure) may play an important role in the flow of groundwater. The estimates presented herein assume the aquifer is unconfined and no barriers to flow exist internal to the Boundary. Improvements to the understanding of groundwater flow may occur as data become available and evaluations are made that re-define geologic controls.

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24 RECOMMENDATIONS

25 26 The following recommendations are suggested to improve estimates of subsurface flow across the Boundary:

- 27 28
- Collect and compile a greater number of GSE measurements proximal to the Ocean and Santa Maria Valley flow zones;
- 29
- Determine the extent that geologic characteristics control groundwater flow;
- 30 31 • Refine understanding of freshwater/saltwater interface and deep groundwater flow within the Paso Robles Formation.
- 32

33 REFERENCES

34 California Department of Water Resources. 2002. (DWR 2002). Water Resources of the Arroyo 35 Grande-Nipomo Mesa Area. September 24, 2002.

To: files Re: Subsurface Flow Date: March 10, 2008 Page: 6 of 6

- 1 Richard C. Slade and Associates. 2000 (Slade 2000). Map of "Base of Fresh Water". March 2000.
- 2 SAIC. 2002. (SAIC 2002). Estimation of Subsurface Inflow for the Hydrologic Inventory for the
- 3 Nipomo Mesa Management Area. June 11, 2002.

D R A F T

¹ Slade 2000, "Base of Fresh Water"
² from Figure 2: Groundwater Elevation Contour Map
⁹ gradient of 0 when subsurface flow is parallel to Boundary
⁴ geometric mean value (DWR 2002)
¹ 1 acre-foot equals 43560 cub

Net Flow

5,050

G)

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Table 4. Water Supply - variables and results

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¹ Pilate 11: Base of Potentially-Water Bearing SedIments (DWR 2002)

² from Figure 2: Groundwater Elevation Contour Map

² gradient of 0 when subsurface flow is parallel to Boundary

⁴ geometric mean value (DWR 200

Net Flow

1,370